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# Physics of the Shearing Process of Saturated Clays

## Physique du Cisaillement des Argiles Saturées

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**SYNOPSIS.** Physical reasons of some mechanical phenomena are proposed for Z.M.bentonite and for some natural claye. Attractive  $/p_A/$  and repulsive  $/p_R/$  long-range particle interaction pressures are calculated and compared with the measured stress values. At high w.c.  $/W > W_p, p_R > -p_A,$  dispersed system/: when  $u=0, \frac{1}{2} \bar{b}'_1 \bar{b}'_3 \approx p_R$ ; when  $\bar{b}'_3=0, p_R \approx -u$  and  $\frac{1}{2} \bar{b}'_1 - \bar{b}'_3 / f \approx p_R$ . At low w.c.  $/W < W_p, -p_A > p_R,$  flocculated system/ interparticle contacts inevitably develop and they carry part of applied pressure: when  $u=0, \frac{1}{2} \bar{b}'_1 \bar{b}'_3 > p_R$ ; when  $\bar{b}'_3=0, p_R > -u$  and  $\frac{1}{2} \bar{b}'_1 - \bar{b}'_3 / f \approx -p_A$ . The condition of  $-p_A = p_R$  is fulfilled at w.c. about  $W_p$  where  $\bar{b}'_3$  was about  $1 \text{ kp/cm}^2$  and half the interparticle distance  $d=35\text{\AA}$ /. During consolidation and/or shearing process average particle thickness may vary as the function of stress applied.

**INTRODUCTION.** The interaction through a thin /rigid/ layer of water is usually assumed to be the cause of true cohesion. Here this water layer is supposed to be of pronounced thickness  $/2d=50\text{\AA}$  to  $100\text{\AA}$ , diffuse layer, non-rigid/ and the long-range interactions /attractive,  $p_A$ , and repulsive,  $p_R/$  are assumed as the reason of the shear strength.

In previous study of this author the comparison of estimated long-range interactions,  $p_R$  and  $p_A$ , with the measured shear strength indicated that either both of them or only one of them may be the cause of clay strength. This problem needed a more detailed research and explanation.

The diffuse layer repulsion pressure,  $p_R$ , was considered in detail by BOLT and MILLER /1955/ and BOLT /1956/. A thorough study of London-van der Waals interaction pressure,  $p_A$ , was performed /STEPKOWSKA, 1970, 1975b/. The water sorption test /WST/ was elaborated, permitting the determination of crystal phase water  $/W_h/p/p_0=0.95; 200^\circ\text{C}/$ , of the external specific surface,  $\bar{S}$ , and of other properties /estimation of CEC, montmorillonite,

M, and kaolinite, K1, content indexes, particle thickness,  $\delta/$ . The eventual interaggregate water,  $W_{mac}$ , may also be estimated and half the interparticle distance,  $d$ , may be obtained for the given w.c. /STEPKOWSKA 1973a and b, 1975b, 1976/:

$$d = \frac{W - W_h - W_{mac}}{\bar{S} \cdot \bar{g}} \quad /1/$$

/here  $\bar{g}=1.0 \text{ g/cm}^3/$  is the density of free liquid water/.

The knowledge of this parameter enables the estimation of  $p_R$  and  $p_A$  for the investigated clay-water system.

Thermodynamic considerations of the mechanical processes in clays indicate that the decreased potential energy,  $dV < 0$ , may be dissipated in form of heat:  $dQ = -dV_R$  in pure tension and  $dQ = -dV_A$  in pure compression /1975a, 1976/.

The theoretical micromechanism of the shearing process presented herein is supported by some check experiments, performed on Z.M.bentonite and on some natural clay samples /M. Ill, partly interstratified, with possible Chl and/or K1 admixtures/. The properties of

the investigated samples are presented in detail elsewhere. Some tests were performed on Sedlec kaolin but their interpretation is not complete as yet.

**ELEMENTS OF CLAY MICROSTRUCTURE.** Crystallite /particle/ is assumed as the smallest elementary component of clay structure. XRD study, WS Test and energetic considerations indicate the following: the crystallite of Z.M.bentonite is composed of several sheets of montmorillonite /plus probably one mica sheet/ of a thickness of  $10\text{\AA}$  each sheet. Number of sheets per particle /Nr/ varies between 5 and 8. On the external surface and in the inter-sheet space there are three-molecular water layers present, of the thickness of  $9\text{\AA}$  to  $10\text{\AA}$  each. The total thickness  $\delta$  of a crystallite is thus  $\sim 100\text{\AA}$  /Nr=5/ to  $\sim 160\text{\AA}$  /Nr=8/. The external specific surface,  $\bar{S}$ , varies between  $\sim 160\text{m}^2/\text{g}$  /Nr=5/ and  $\sim 100\text{m}^2/\text{g}$  /Nr=8/.

In certain natural clay /H.K./ containing little or no quartz,  $\delta$  measured by XRD and WS Test was  $200\text{\AA}$  to  $350\text{\AA}$  and the corresponding values of  $\bar{S}$  were 65 to  $25\text{m}^2/\text{g}$ . In clays from some other localities the particle thickness  $\delta$  measured was  $150\text{\AA}$  to  $300\text{\AA}$  /D.clay/,  $200\text{\AA}$  to  $220\text{\AA}$  /Bydgoszcz clay/,  $700\text{\AA}$  to  $800\text{\AA}$  /B.clay, from WST/.

Clay crystallites are arranged parallel to each other /due to diffuse layer repulsion,  $p_R$ /, forming domains or tactoids. The mutual distance depends on w.c. and it may vary /Z.M.bentonite/ between  $\sim 140\text{\AA}$  at liquid limit and  $\sim 70\text{\AA}$  at plastic limit,  $w_p$ , around which value the repulsion equals attraction. In the case of absence of interparticle links and of presence of bivalent exchangeable cations / $\delta \sim 100\text{\AA}$  to  $140\text{\AA}$ /:

$$\text{at } w = w_p \text{ and } d = 70\text{\AA}, \quad p_R = -p_A \quad /2/$$

Below this w.c. /attraction exceeding repulsion/ interparticle bonds develop either due to ionic lattice attraction between sheets curled at the crystallite edges /montmorillonite/ or due to crystallite reorientation and edge-to-face contacts probable in kaolinite. These bonds impede further decrease in interparticle distance /with load increase/ as the

bending strength of crystallites is much higher than  $p_R$ . Unbending of crystallites after unloading may influence the suction measured / $u < 0$ /.

Domains /tactoids/ in random mutual orientation form aggregates /peds/, of the size which is measured as grain size. In natural clays interaggregate water,  $w_{mac}$ , may occur, in Z.M.bentonite /N.C./ EM-study and calculation result indicate that  $w_{mac} = 0$ .

For a group of preconsolidated natural clay samples from a given locality this value can be estimated from WS Test, i.e. from the intercept of the regression line /transformed Eq. 1/:

$$w - w_h = w_{mac} + d_{av} \bar{S} \quad /3/$$

In highly O.C. clay /e.g. due to glacier/,  $d$  may be assumed as close to an average value,  $d_{av}$ . The regression line calculated for 11 samples of B.clay /mainly from the depth exceeding 20 m/ was:

$$w_f - w_h = 4.64 + 0.285 \bar{S} / \%, \quad r_{xy} = 0.9795 \quad /3a/$$

indicating  $w_{mac} = 4.64\%$  and  $d_{av} = 28.5\text{\AA}$ .

**MICROMECHANISM OF THE SHEARING PROCESS.** Isotropic consolidation of Z.M.bentonite from w.c. near to the liquid limit  $w_o \approx 100\% / \epsilon_o = 0.25$  to  $5.0\text{kp}/\text{cm}^2$  was performed in Norwegian type triaxial test equipment. Samples were sheared after unloading to  $\epsilon_3 = 0$  with pore water pressure,  $u$ , measurement /Plexi null device/ at a constant strain rate of  $1.6\text{ mm}/\text{hour}$ . The lines in Fig. 1 represent the calculated values of  $p_R$  and  $p_A$  for concentration range of  $10^{-3}$  to  $10^{-4}\text{ n}$  and particle thickness range,  $\delta = 100\text{\AA}$  to  $140\text{\AA}$  /most probable values in investigated bentonite/. The values of  $p_R$  are approximately valid also for natural clays; usually their  $\delta$  and  $-p_A$  values are higher.

I. "Dispersed" structure:  $p_R > -p_A$ ,  $w > w_p$ ,  $d > 35\text{\AA}$ ,  $\epsilon_c \leq 1\text{kp}/\text{cm}^2$ . If the initial w.c. is high enough,  $\epsilon_c$  is carried mainly by  $p_R$  and after termination of the consolidation process /pure compression/:

$$\text{N.C. } /u=0/ \quad \frac{1}{2} \epsilon_c = p_R \quad /4/$$

The decreased potential energy of attraction is dissipated in form of heat /1975a, 1976/:

$$-dV_A = dQ_A \quad /5/$$

After undrained unloading /  $\epsilon_3=0$  /:

$$p_R = -u \quad /6/$$

Experimentally measured suction values /Fig.1/ are close to the calculated  $p_R$  values. Small differences may be due to  $p_A$  impeding reversible unloading, to unbending of interparticle links and to experimental and/or estimation error.

The shearing process may be presented here as pure compression / $dd < 0$ /. Work is performed only when decreasing the interparticle distance

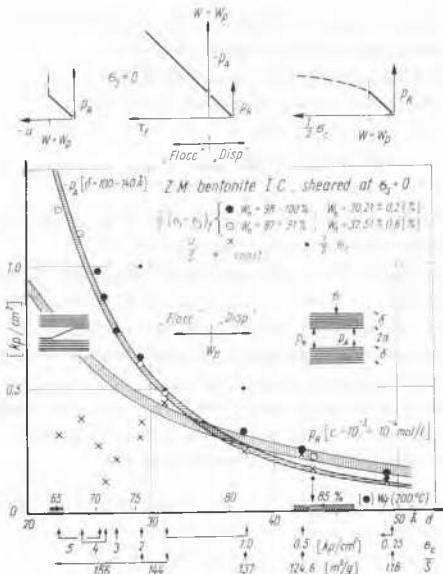


Fig 1

no work is done when the particles are displaced parallel to each other /the potential gradient in this direction equals zero/. Work of external forces:

$$dW = \sigma_d d \epsilon_d \quad /7/$$

causes a corresponding change in potential energy of repulsion:

$$-dV_R = p_R dd \quad /8/$$

where  $d \epsilon_d = 2 dd \neq 0$  /9/

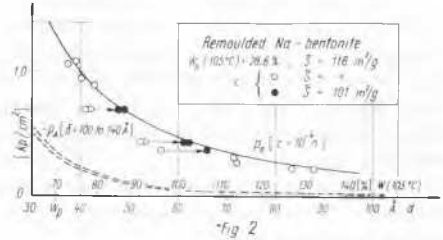
When particle orientation is random, when /7/ equals /8/ and Equation /5/ is true, this leads to the general expression of Eg. /4/.

In unconfined compression / $\epsilon_3=0$  /:

$$\frac{1}{2} \sigma_1 - \sigma_3 / f = \frac{1}{2} \sigma_1 / f = p_R \quad /10/$$

This was measured in Z.M.bentonite /Fig.1/ at  $\sigma_c < 1.0 \text{ kP/cm}^2$ . Sheared samples failed mainly by bulging.

Remoulded Na-bentonite and /Na+Ca/-bentonite, indicating high  $p_R$  values, behaved in quick triaxial compression in accordance with the above presentation /Fig. 2/.

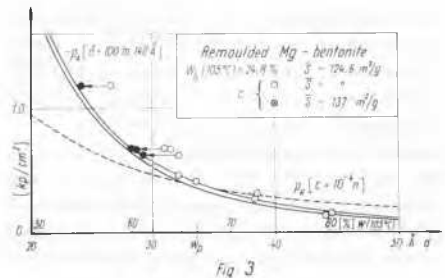


II. "Flocculated" structure:  $-p_A > p_R$ ,  $W < W_p$ ,  $d < 35 \text{ \AA}$ ,  $\sigma_c > 2 \text{ kP/cm}^2$ . Here a part of the consolidation pressure is carried by the soil skeleton and only the rest of it is carried by  $p_R$ . Thus at N.C./ $u=0$ /  $\frac{1}{2} \sigma_c \gg p_R$  /11/

On unloading in undrained condition:

$$\epsilon_3=0 \quad -u \leq p_R \quad /12/$$

The swelling of the sample in drained condition is limited both by the attractive pressure



$$-dV_R = dQ_R \quad /13/$$

and partly due to crystallite deformation.

Work of external forces /7/ causes change in the potential energy of attraction:

$$-dV_A = p_A dd \quad /14/$$

Assuming /7/=14/ and /9/

$$\frac{1}{2} \epsilon_1 - \epsilon_3 / f = -p_A \quad /15/$$

This was approximately observed in Z.M.bentonite samples in unconfined compression  $\epsilon_3=0$ , see Fig.1, and in triaxial compression of N.C. and O.C. samples /results will be presented elsewhere/. At failure the sheared samples mostly indicated a distinct shear plane.

Remoulded Mg-, Ca-, Fe-, Al, and H- bentonites behaved in accordance with this presentation /see Fig. 3/.

Perfect correlation between the theoretically calculated forces and the measured shear strength was obtained, assuming a stepwise increase in external specific surface,  $\bar{S}$ , /decrease in  $Nr$ / with the increase in  $\epsilon_c$  and/or shear strength. Though the ionic lattice attraction has a very high component perpendicular to the sheet face, its component parallel to the sheet face is zero and creep phenomenon is probable. The values of  $Nr$  and  $\bar{S}$  assumed in calculation were equal to values determined by XRD /particle thickness/ for samples oriented from a slurry and for powder samples /WS Test gave a good check/. The values assumed were:

$$\begin{array}{lll} \epsilon_c < 1 \text{ kp/cm}^2 & Nr=6,5 \text{ or } 7 & \bar{S}=124.6 \text{ or } 118 \text{ m}^2/\text{g} \\ =1 \text{ kp/cm}^2 & =5,5 \text{ or } 6 & =144 \text{ or } 137 \text{ m}^2/\text{g} \\ =2 \text{ kp/cm}^2 & =5 \text{ or } 5,5 & =156 \text{ or } 144 \text{ m}^2/\text{g} \\ > 2 \text{ kp/cm}^2 & =5 & =156 \text{ m}^2/\text{g} \end{array}$$

A unique relation was obtained between the shear strength and the interparticle distance  $2d$  for all the investigated Z.M.bentonite samples /N.C., O.C., remoulded and statically compacted/. Similar results were obtained for a series of stiff H.K. clay samples, preconsolidated by a glacier /1975b/. Interparticle distance was estimated by WS Test. The average value of calculated therefrom  $-p_A$  was  $0.74 \text{ kp/cm}^2$ . The average value of shear strength measured was  $0.71 \text{ kp/cm}^2$ .

The following regression lines for H.K. clay samples were calculated:

$$\begin{array}{l} 1/ \frac{1}{2} \epsilon_1 - \epsilon_3 / f = -0.0522 d_f + 2.411 \quad / \text{kp/cm}^2 / \\ r_{xy} = -0.5781 \quad /16/ \end{array}$$

$$\begin{array}{l} 2/ \frac{1}{2} \epsilon_1 - \epsilon_3 / f = 1.043 \quad [-p_A / \delta = 300 \text{ R} /] + 0.004 \\ / \text{kp/cm}^2 / \quad r_{xy} = 0.7980 \quad /17/ \end{array}$$

Numerical value of correlation coefficient of Eq./17/, as compared to that of Eq./16/, indicates that  $p_A$  is responsible for the shear strength.

B. clay from the depth of 2 to 11.5m, sheared at  $\epsilon_3=1$  to  $2 \text{ kp/cm}^2$ , indicated shear strength of  $1.52$  to  $2.10 \text{ kp/cm}^2$ , whereas the average estimated theoretical value was  $-p_A=1.54 \text{ kp/cm}^2$  / $\delta=700 \text{ R}$ ,  $d=28,5 \text{ R}$ , see Eq. 3a/.

CONCLUSIONS. In saturated swelling /montmorillonitic/ clays, indicating no cementation and no non-clay mineral admixtures:

1. The reason of the shear strength is:
  - a. London-ven der Waals attraction  $p_A$   
if  $-p_A > p_R$  and  $W < W_p$
  - b. diffuse layer repulsion  $p_R$   
if  $p_R > -p_A$  and  $W > W_p$
2. Plastic limit,  $W_p$ , is close to w.c. where attraction equals repulsion.
3. Increase in stress applied may cause a decrease in average particle thickness and an increase in average shear strength /shear plane passing through crystallite parallel to its surface/.

These conclusions were drawn for Z.M.bentonite and Polish natural clays and they should be checked for other types of soils.

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