Analysis of the impact of climate change on the landslides along the mountain highway in central Taiwan

ABSTRACT: Climate change caused by global warming affects Taiwan significantly for the past decade. The increasing frequency of extreme rainfall events, in which concentrated and intensive rainfalls generally cause geohazards including landslides. Considering the existence of various types of landslides and the protection targets, this study aims to analyse the landslide susceptibility along the Nantou County Road # 89 in Central Taiwan. This study employs rainfall frequency analysis and the atmospheric general circulation model (AGCM) downscaling estimation to obtain the temporal rainfall distributions, then the landslide susceptibility was estimated for the study area. Based on the Li-DAR and the borehole data, the temporal behavior and the complex mechanism of large scale deep-seated landslides were also analyzed.

KEYWORDS: landslide hazard, climate change, mountain highway, landslide susceptibility

1 INTRODUCTION

The climate change affects Taiwan significantly by an increasing frequency of extreme rainfall events, in which induced large scale landslides. Considering the existence of various types of large scale landslides and the protection targets, this study aims to analyze the landslide susceptibility along the Nantou County Road # 89 of Taiwan (see Fig.1). For the predictive analysis of landslide susceptibility, this study employed AGCM downscaling estimation. For the adopted large scale landslides (see Fig. 2), based on the information from boreholes, the temporal behavior and the complex mechanism of large scale landslides were analyzed in local scale. Based on the results, the pros and cons of the analysis in both scale were discussed, which could be applied for the risk assessment and management.

Figure 1. Geology of the study area

Figure 2. Large landslide along the Nantou County Road # 89

2 METHODOLOGIES

Landslide inventories, SPOT satellite images, borehole data were collected in the study area, including 5 major deep-seated landslides along the County Road #89. In addition, mobile laser ranging was performed along the road to obtain the 0.1 LiDAR data. The obtain information and data were used for the shallow and deep-seated landslides in different scale.

2.1 Landslides interpretation

This study adopts the NDVI-slope angle criterion, in which the normalized differential vegetation index (NDVI) is from satellite images and the slope angle is from digital elevation model (DEM). And improving the accuracy of landslide identification in shadow areas with different screening indexes,
including brightness (BRI, Hsieh et al 2011), greenness (GI, Liu et al 2012; Lin et al 2013), and vegetation mask (Beumier and Idrissa 2014). The compare of landslides interpretation in different indexes as Table 1. A1 is the number of landslide cells interpreted as landslide, A3 is the number of landslide cells not interpreted as landslide. A2 is the number of non-landslide cells interpreted as landslide, A4 is the number of non-landslide cells interpreted as non-landslide.

**Table 1. Comparison of the Criteria for Automatic Landslide Interpretation.**

<table>
<thead>
<tr>
<th>Criterion</th>
<th>Accuracy of Landslide Cells A1/(A1+A3)</th>
<th>Accuracy of Non-landslide Cells A4/(A2+A4)</th>
<th>Accuracy of Total Cells A1+A0/(A1+A2+A3+A4)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Slope=20%, NDVI=0, BRI=40</td>
<td>6.60</td>
<td>99.63</td>
<td>98.49</td>
</tr>
<tr>
<td>Slope=20%, NDVI=0, GI=0.25</td>
<td>7.11</td>
<td>99.62</td>
<td>98.49</td>
</tr>
<tr>
<td>Slope=50%, NDVI=0.2</td>
<td>21.02</td>
<td>97.79</td>
<td>96.86</td>
</tr>
<tr>
<td>Slope=20%, NDVI=0.2, BRI=40</td>
<td>18.81</td>
<td>97.87</td>
<td>96.90</td>
</tr>
<tr>
<td>Slope=20%, NDVI=0.2, GI=0.25</td>
<td>20.40</td>
<td>97.92</td>
<td>97.00</td>
</tr>
<tr>
<td>Slope=20%, NDVI=0.2, BRI=60</td>
<td>7.04</td>
<td>98.46</td>
<td>97.34</td>
</tr>
</tbody>
</table>

2.2 Landslide-rainfall index (Id)

For a specific typhoon event, by the overlapping function of GIS, the accumulated rainfall and rainfall intensity data at the landslide locations can be extracted and plotted in graph of accumulated rainfall and rainfall intensity (see Fig. 3). The landslide-rainfall index (Id) is defined by the distances d1 and d2 from the unknown point to the linear thresholds as d2/(d1+d2) (Shou et al, 2015).

![Figure 3. The landslide-rainfall index (Id) is defined as d2/(d1+d2).](image)

It ranges between 0 and 1. As Id approaches 1, the slope becomes increasingly susceptible to rainfall-induced landslide. On the contrary, as the point of the rainfall of potential landslide approaches the lower threshold, or as Id approaches 0, the slope becomes less susceptible to landslide.

2.3 Landslide susceptibility models

This study adopts the In Logistic Regression Method for the landslide susceptibility analysis. Its performance was compared for the analyses of 2004 Mindulle, 2009 Morakot, and 2012 Saola. Based on the training samples, which comprised a group of data points or data locations, categorized as landslide and non-landslide. The data layer of each factor was then placed upon the landslide and non-landslide layers, and the correlation between each factor and landslides was used to conduct binary logistic regression. For the susceptibility model obtained by logistic regression, this study employed the receiver operating characteristic (ROC) curve (Swets 1988), in which the area under the curve (AUC) of the ROC curve was used to evaluate the prediction accuracy. Generally, the larger the AUC values the better. As the area approaches 0.5, the result may not necessarily be superior to that of a random selection. AUC values of less than 0.5 are not worth employing.

3 LANDSLIDE SUSCEPTIBILITY ANALYSIS

This study adopts the Logistic Regression Method for the landslide susceptibility analysis. The model based on 2004 Mindulle can be expressed as

\[ P = 0.663F1 + 0.055F2 - 0.317F3 - 0.062F4 - 0.301F5 - 0.216F6 - 0.093F7 + 0.081F8 - 2.451F9 + 0.295 \]

where P is the logistic function, F1 is the slope angle, F2 is the elevation, F3 is the aspect, F4 is the distance to fault, F5 is the distance to river, F6 is the distance to road, F7 is the dip slope index (Ids), F8 is the landslide-rainfall index (Id), and F9 is the normalized differential vegetation index (NDVI). Eq. (1) can be used to calculate the landslide susceptibility based on the predicted rainfalls, including various extreme weather scenarios as below.

![Figure 4. The landslide susceptibility along Nantou County Road #89 estimated by 2004 Mindulle model](image)
4 ANALYSIS OF MAJOR LARGE SCALE LANDSLIDES

For the adopted large scale landslides (see Fig. 2), based on the information from boreholes, the temporal behavior and the complex mechanism of large scale landslides were analyzed in local scale. Fig. 6 shows locations of the 5 studied major profile. In this study, the limit equilibrium analyses software Slide 6.0 (Rocscience 2015) was applied.

In the analyses, three types of geo-material were assumed, i.e., highly weathered slate in the boundary layer, medium weathered slate in the middle layer, and fresh slate in the bottom. Their material properties were determined based on the report of exploration (Xie-Sheng Engineering Consulting, 2014). For a better comparison, this study considered three different groundwater conditions, i.e., dry (no groundwater), 1/2 ground water level, and full ground water level. The results of equilibrium analyses were summarized in Table 3.
5 CONCLUSION

In this study, rainfall frequency analysis and the atmospheric general circulation model (AGCM) downscaling estimation were applied to understand the temporal rainfall trends and distributions in the study area. The susceptibility analysis in catchment scale and local scale were performed for the hazard assessment of the mountain highway, i.e., Nantou County Road #89 in Central Taiwan. The hazard of the major landslides can be ranked to prioritize the hazard mitigation. It is worth noting that the results of local scale analysis also suggest a similar hazard ranking of these landslides, i.e. the sites 15k+850 and 12k+300 are the most dangerous.

6 ACKNOWLEDGEMENTS

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7 REFERENCES


